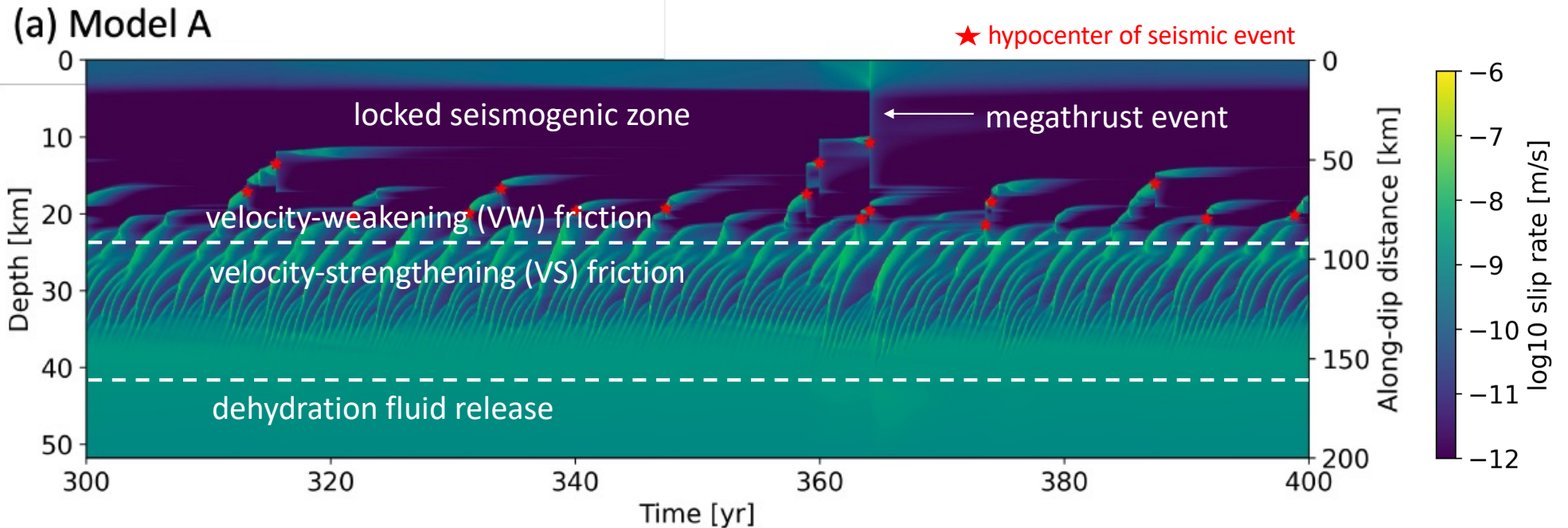


Fault-Valve Instability: A Mechanism for Slow Slip Events

Eric M. Dunham¹, So Ozawa^{1,2}, Yuyun Yang^{1,3}

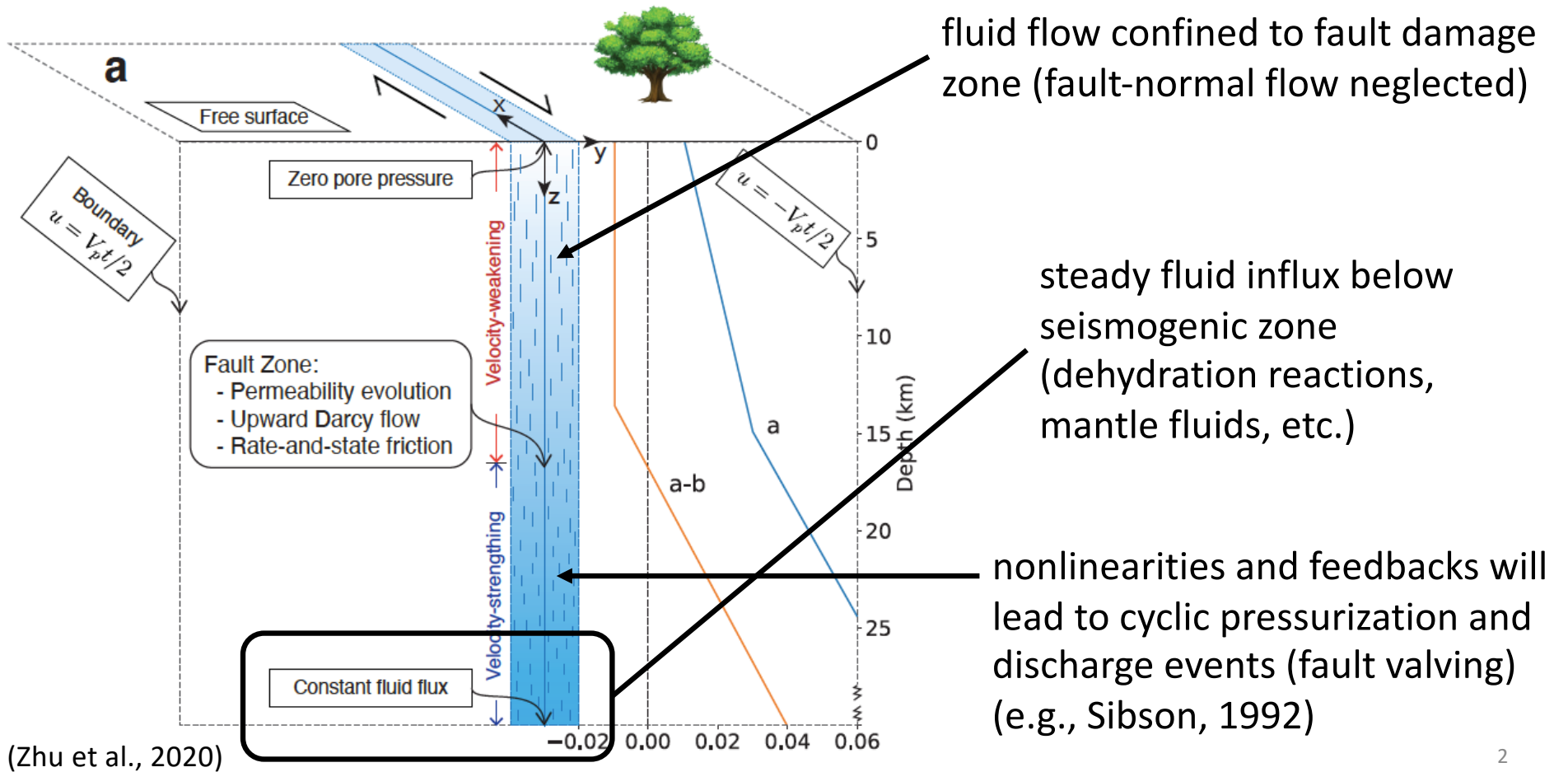
1 Stanford University, 2 now Earthquake Research Institute, Japan, 3 now Chinese University Hong Kong

June 24, 2024



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Fault zone fluid migration and pore pressure evolution models



Idealized model for permeability evolution

elastic dependence on
effective stress:

$$k = k^* e^{-(\sigma-p)/\sigma^*}$$

stress sensitivity
parameter ~ 10 MPa

minimally parametrized (but ad hoc) evolution equation for permeability:

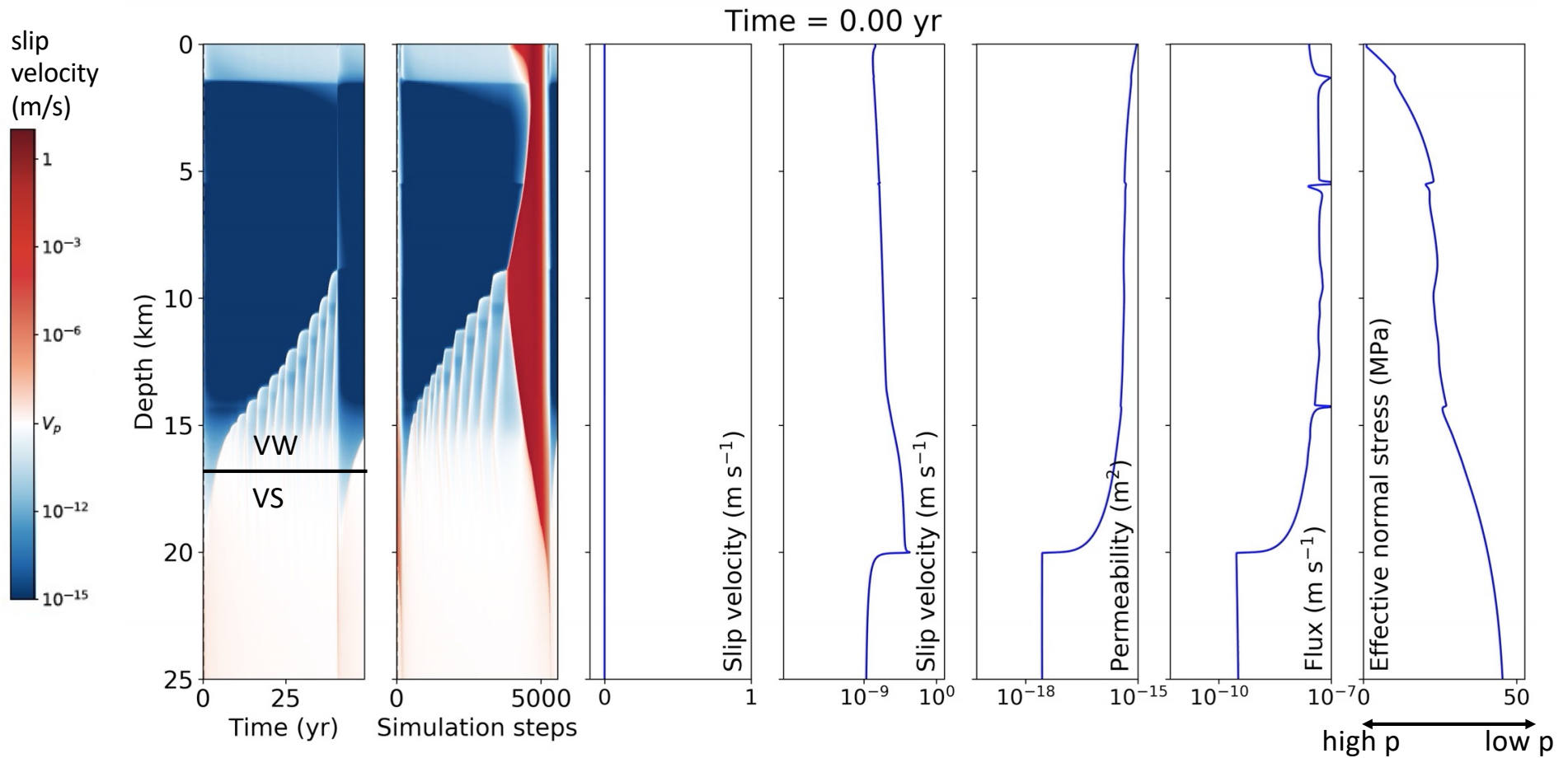
$$\frac{\partial k^*}{\partial t} = -\frac{V}{L} (k^* - k_{\max}) - \frac{1}{T} (k^* - k_{\min})$$

permeability
increase with slip

permeability
decrease from
healing/sealing

→ constant T and L in Zhu et al. (2020), later depth-dependent

Our original fault valving cycle simulations



(Zhu et al., 2020)

Motivation for additional study

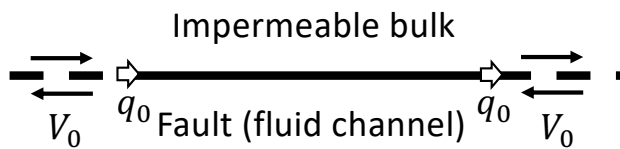
- Which processes and conditions are required for instability?
- Quantitative expressions for recurrence interval, slip pulse length, etc.
- Sensitivity to model assumptions (e.g., permeability evolution model)

Study with spatially uniform properties in homogeneous whole-space

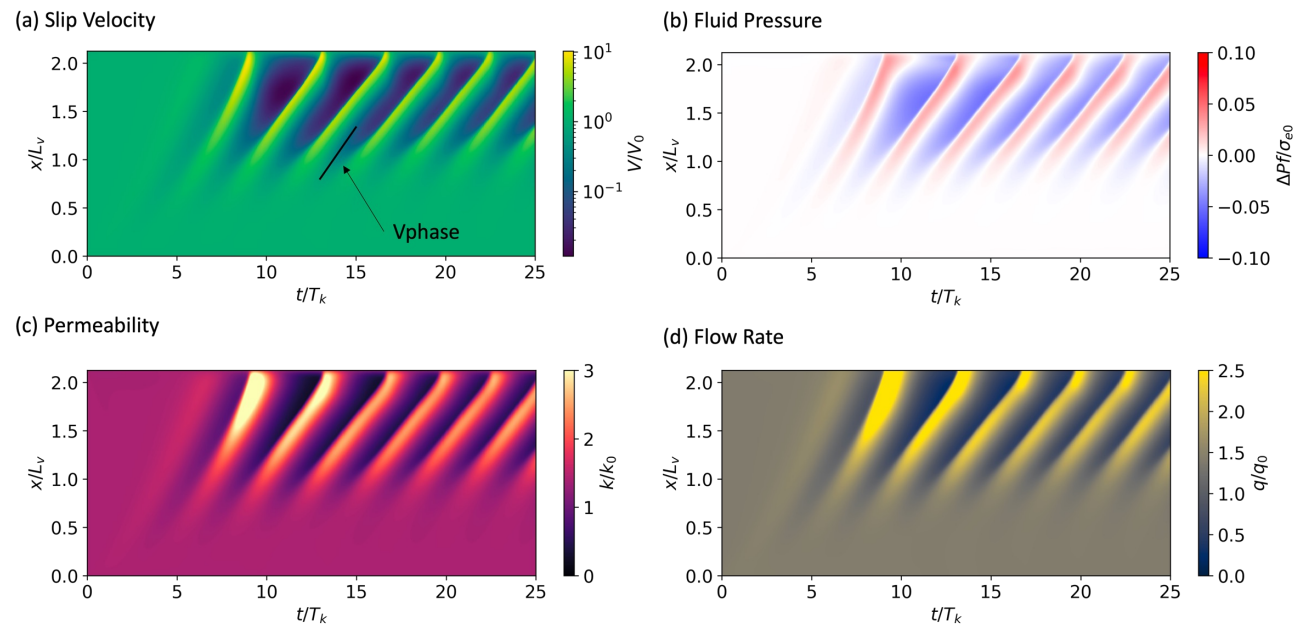
- Linear stability analysis for perturbations about steady sliding (V_0) and steady fluid flow (q_0)
- Nonlinear simulations

Fluid flow destabilizes steady sliding, even with velocity-strengthening friction

(example below)



- fluid-driven aseismic slip fronts propagate in direction of fluid flow
- recurrence interval (and slip/event) controlled by permeability evolution timescale



(Ozawa, Yang, & Dunham, 2024, in review)

Fault valve instability mechanism – fluid flow

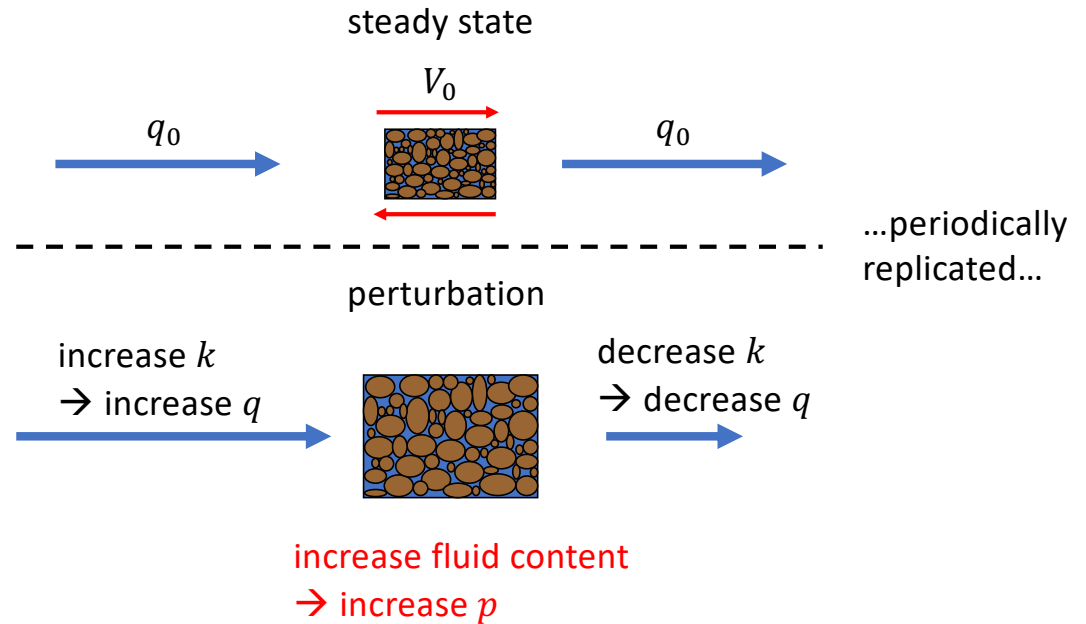
Linearization of pressure diffusion:

$$\beta\phi \frac{\partial p}{\partial t} - \frac{\partial}{\partial x} \left(\frac{k}{\eta} \frac{\partial p}{\partial x} \right) = 0.$$

$$\beta\phi \frac{\partial(p_0 + p')}{\partial t} - \frac{\partial}{\partial x} \left(\frac{k_0 + k'}{\eta} \frac{\partial(p_0 + p')}{\partial x} \right) = 0.$$

$$\beta\phi \frac{\partial p'}{\partial t} - \frac{k_0}{\eta} \frac{\partial^2 p'}{\partial x^2} + \frac{q_0}{k_0} \frac{\partial k'}{\partial x} = 0,$$

diffusion stabilizing



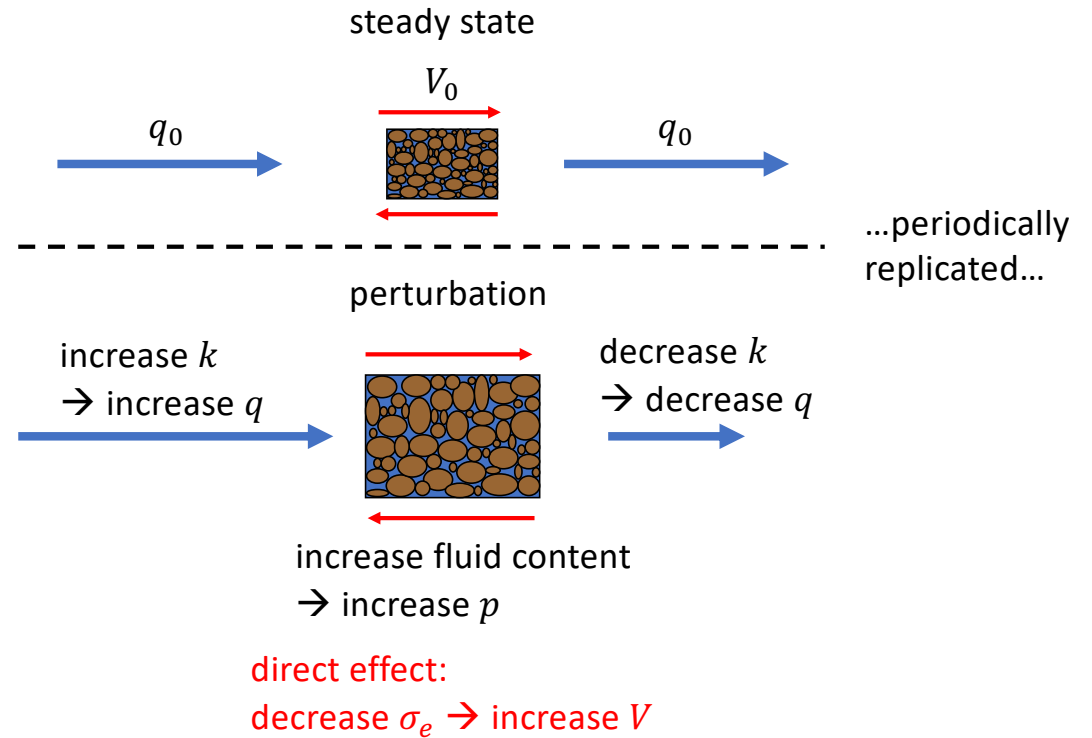
Fault valve instability mechanism – frictional sliding

Linearization of fault strength:

$$\frac{d\tau}{dt} = \frac{a\sigma_0}{V_0} \frac{dV}{dt} + f_0 \frac{d\sigma_e}{dt} - \frac{V_0}{d_c} [\tau - \tau_{ss}(\sigma_e, V)],$$

$$\tau_{ss}(\sigma_e, V) = \tau_0 + f_0(\sigma_e - \sigma_0) + (a - b)\sigma_0 \frac{V - V_0}{V_0}.$$

elasticity stabilizing



Fault valve instability mechanism – frictional sliding

Linearization of fault strength:

$$\frac{d\tau}{dt} = \frac{a\sigma_0}{V_0} \frac{dV}{dt} + f_0 \frac{d\sigma_e}{dt} - \frac{V_0}{d_c} [\tau - \tau_{ss}(\sigma_e, V)],$$

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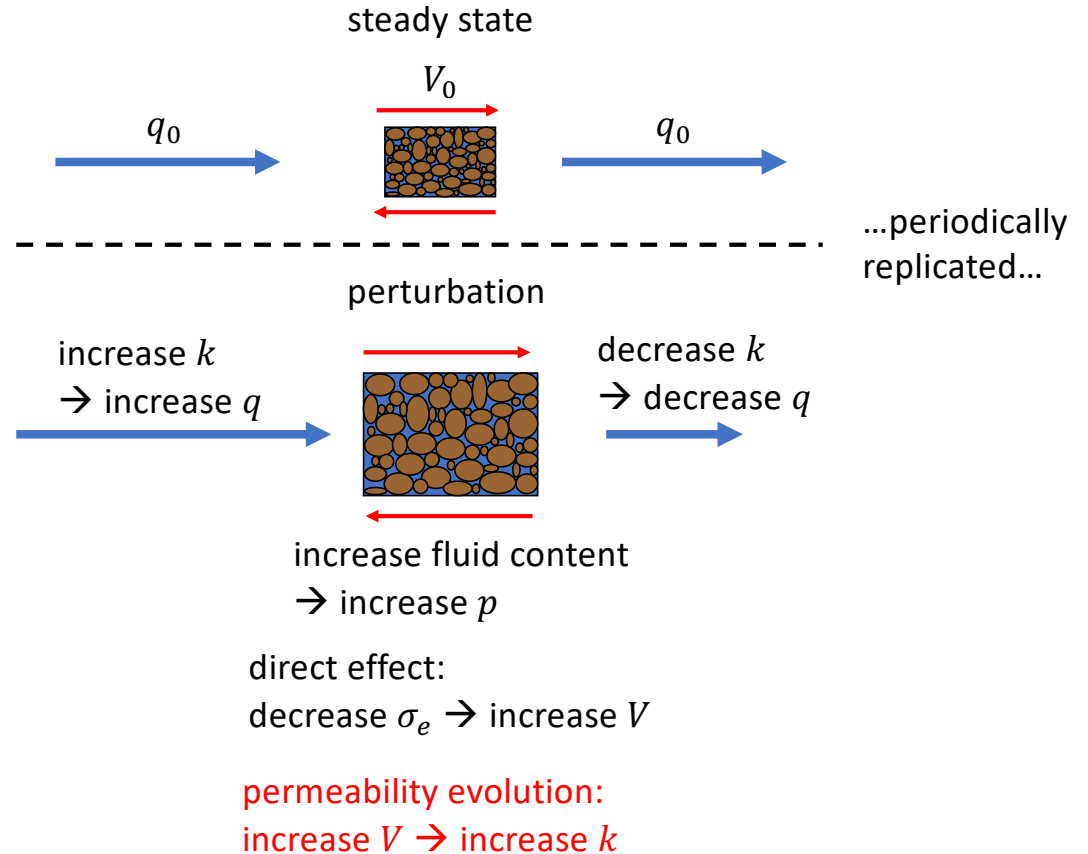
elasticity stabilizing

Linearized permeability evolution:

$$\frac{dk}{dt} = -\frac{k_0}{\sigma^*} \frac{d\sigma_e}{dt} - \frac{1}{T_k} [k - k_{ss}^{lin}(V, \sigma_e)],$$

$$k_{ss}^{lin}(V, \sigma_e) = k_0 - k_0 \frac{\sigma_e - \sigma_0}{\sigma^*} + \Delta k \frac{V - V_0}{V_0}$$

but $k = k_{ss}(V)$ gives *wrong phase relation, no unstable feedback*



Fault valve instability mechanism – frictional sliding

Linearization of fault strength:

$$\frac{d\tau}{dt} = \frac{a\sigma_0}{V_0} \frac{dV}{dt} + f_0 \frac{d\sigma_e}{dt} - \frac{V_0}{d_c} [\tau - \tau_{ss}(\sigma_e, V)],$$

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elasticity stabilizing

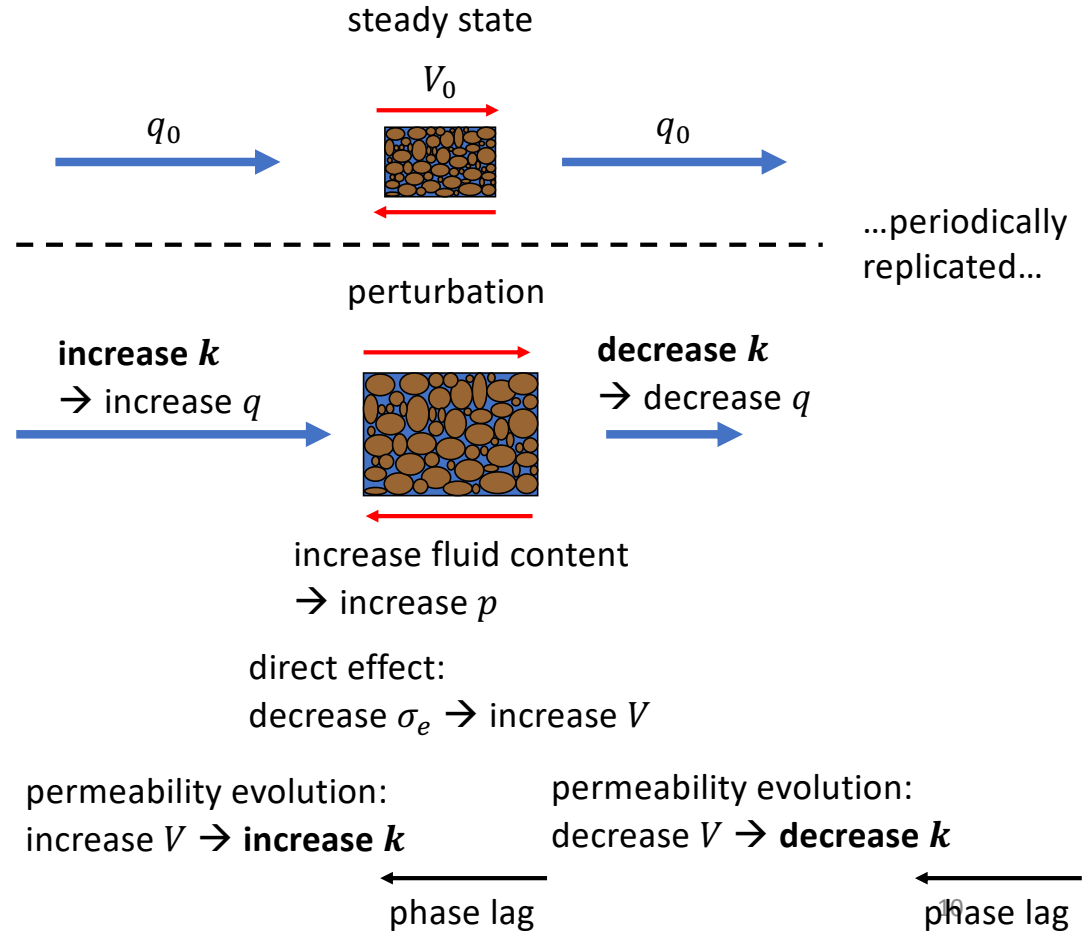
Linearized permeability evolution:

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$$k_{ss}^{lin}(V, \sigma_e) = k_0 - k_0 \frac{\sigma_e - \sigma_0}{\sigma^*} + \Delta k \frac{V - V_0}{V_0}$$

adding evolution of k toward $k_{ss}(V)$
lags phase → **unstable feedback!**

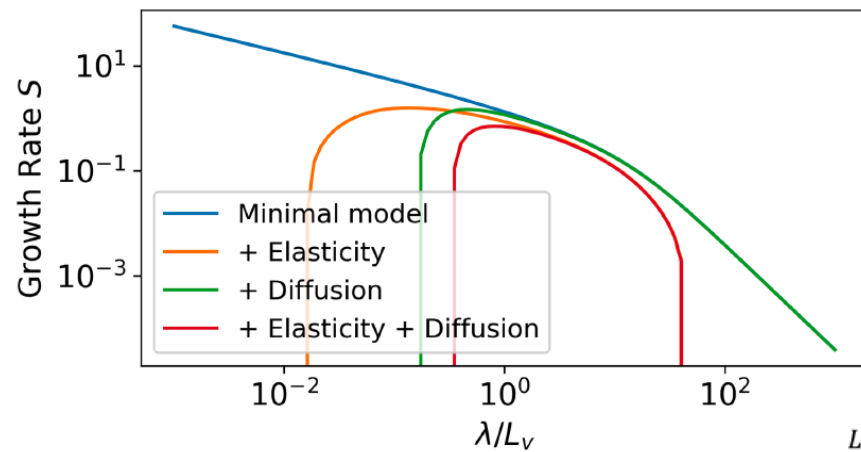
(Ozawa, Yang, & Dunham, 2024, in review)



Linear stability analysis predicts wavelength of max growth

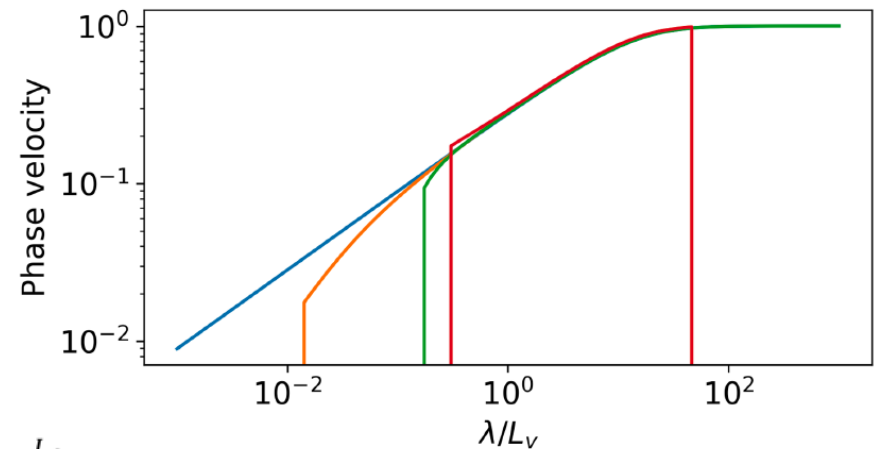
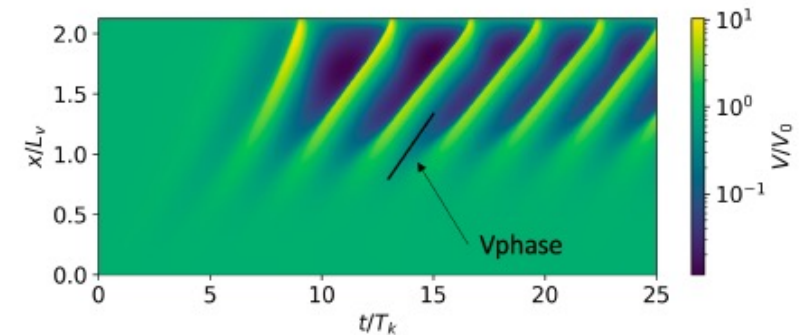
Wavelengths $\sim L_v$ and timescales $\sim T_k$
selected in nonlinear simulations

$$L_v = \frac{f_0 q_0 \Delta k T_k}{k_0 \beta \phi a \sigma_0} \quad (\text{fault valve length scale})$$



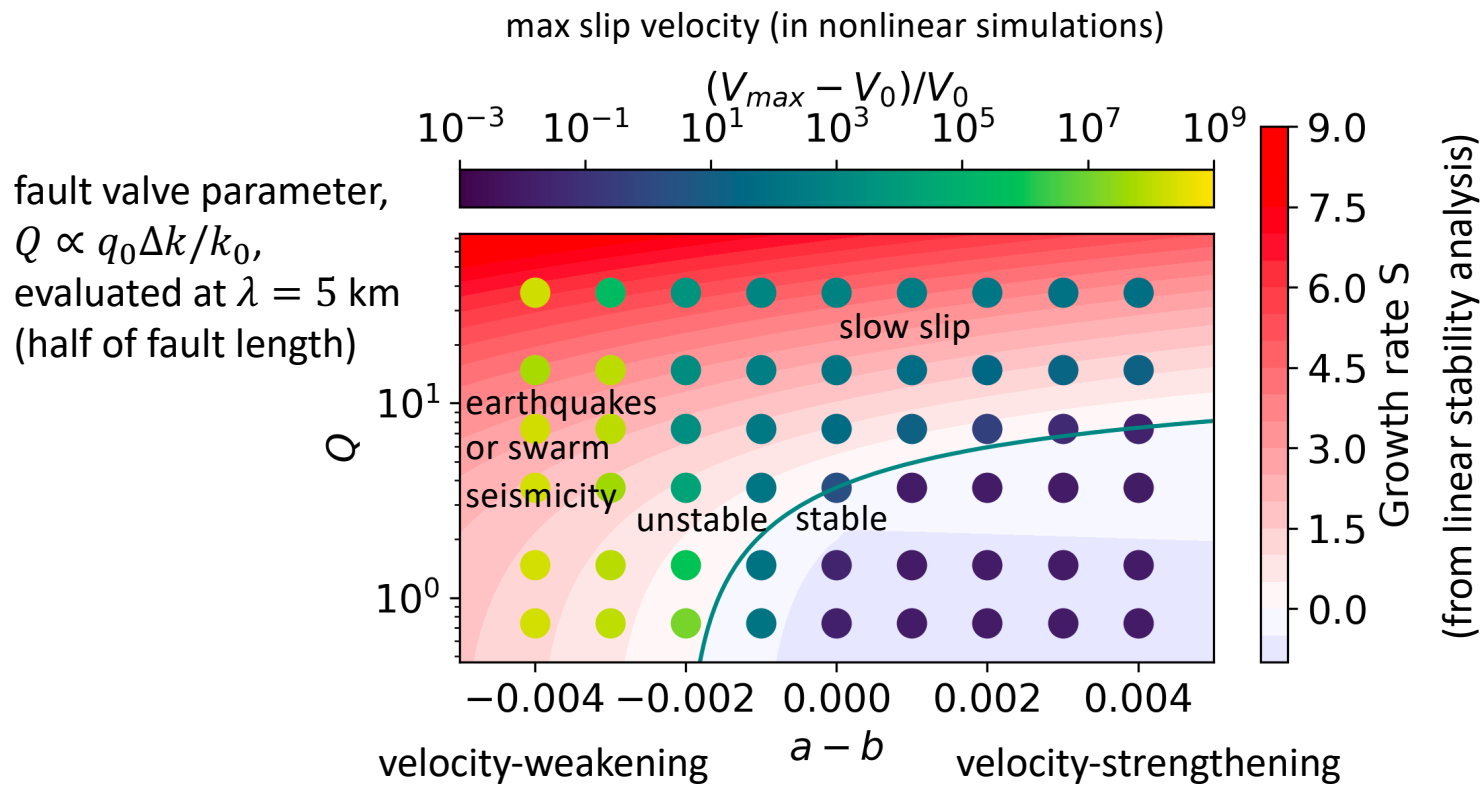
$$\frac{L_e}{L_v} = 0.27, \frac{L_d}{L_v} = 0.004$$

(a) Slip Velocity



(Ozawa, Yang, & Dunham, 2024, in review)

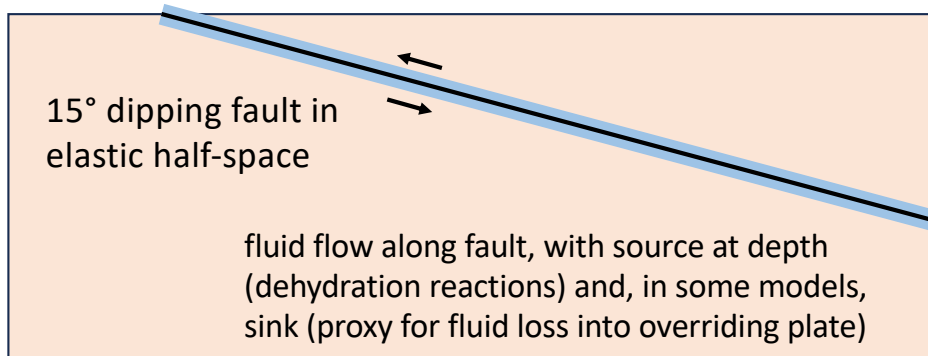
Transition between fault valve instability and classical rate-state frictional instability for VW friction



Subduction zone simulations

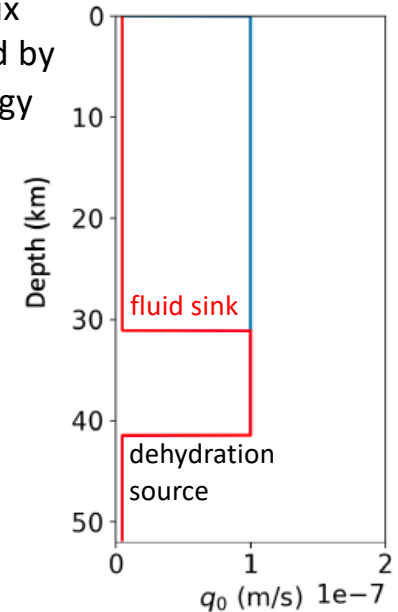
Integrate $q = \frac{k}{\eta} \left(\frac{\partial p}{\partial z} - \rho g \right) \rightarrow$ steady state $p(z)$ and $k(z)$ distributions, then unsteady...

(4 models A-D, different colors in plots to follow)



+ aging law rate-state friction, transition from VW to VS around 22 km depth

fluid flux inspired by petrology

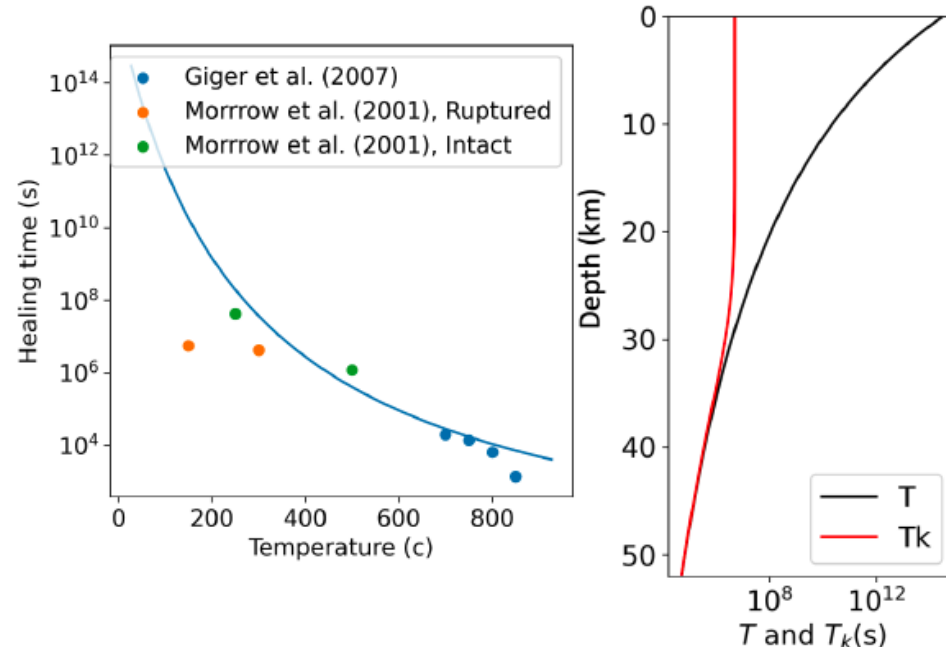


Subduction zone simulations

Integrate $q = \frac{k}{\eta} \left(\frac{\partial p}{\partial z} - \rho g \right) \rightarrow$ steady state $p(z)$ and $k(z)$ distributions, then unsteady...

(4 models A-D, different colors in plots to follow)

temperature and hence
depth-dependent healing time



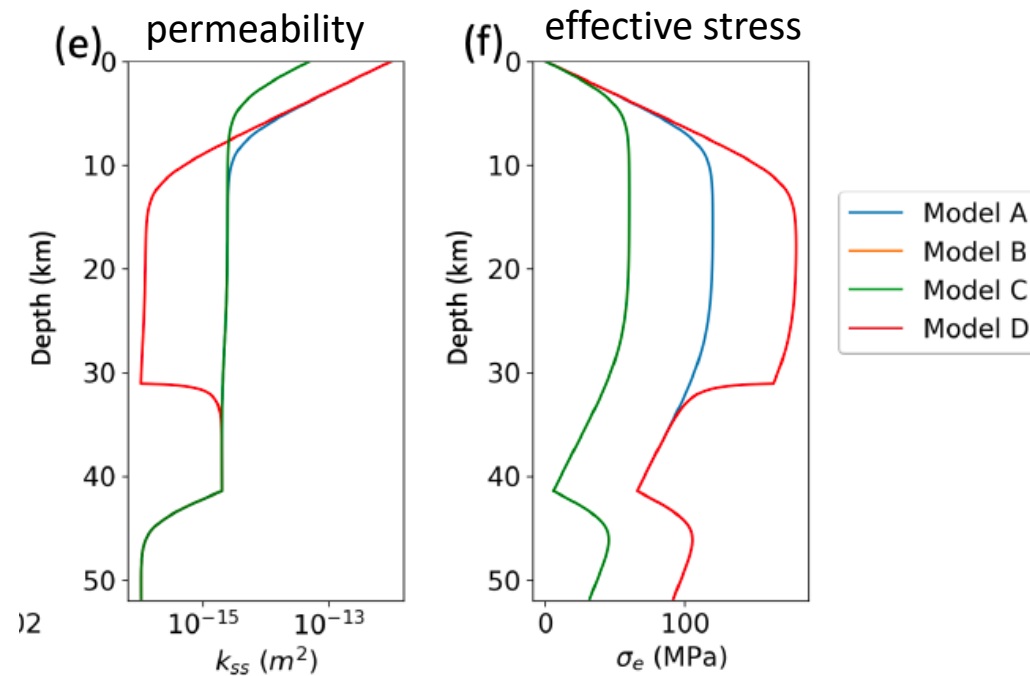
(Ozawa, Yang, & Dunham, 2024, in review)

Subduction zone simulations

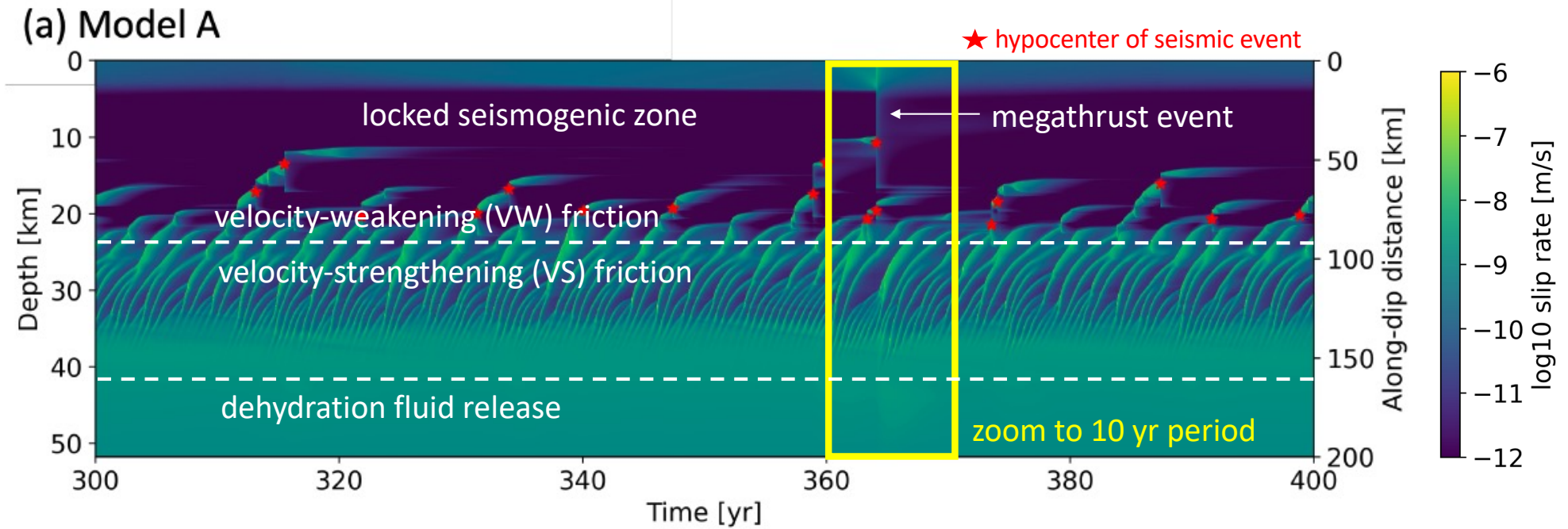
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(4 models A-D, different colors in plots to follow)

steady state
solutions



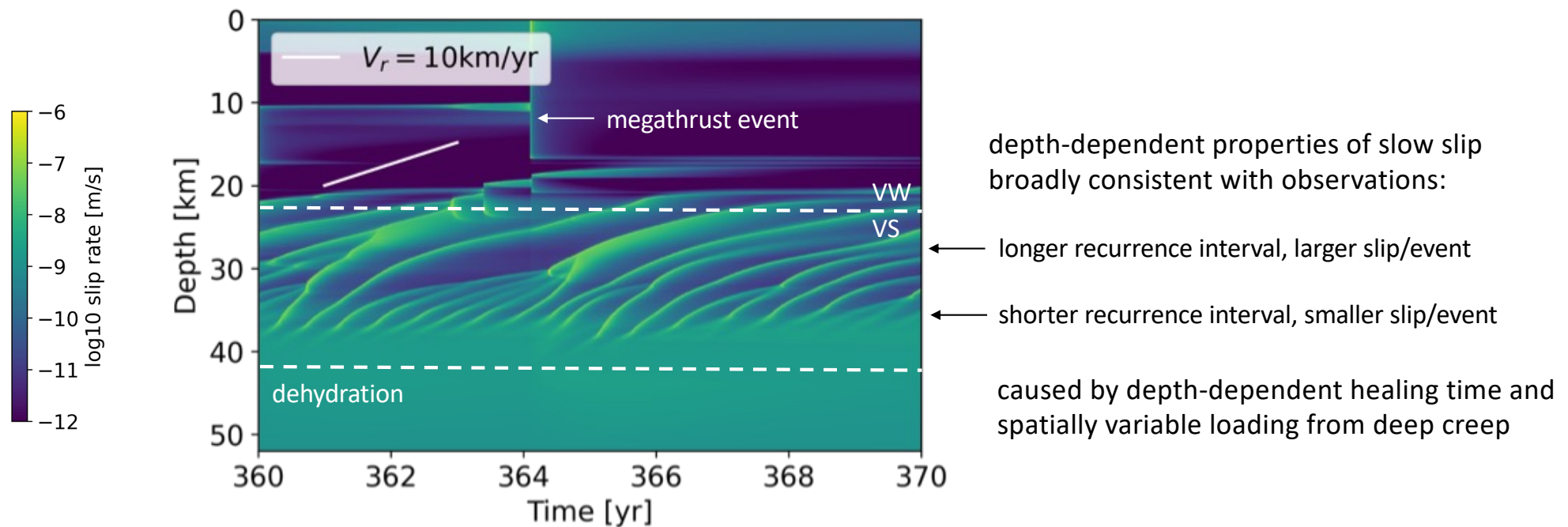
Reference subduction model



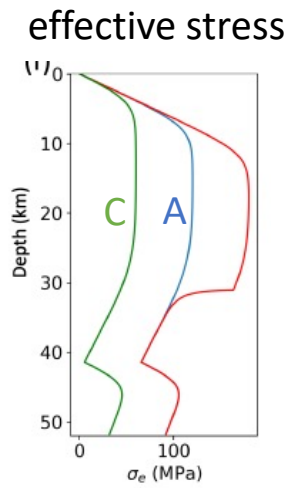
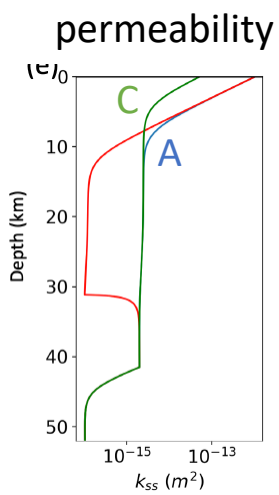
(Ozawa, Yang, & Dunham, 2024, in review)

Reference subduction model

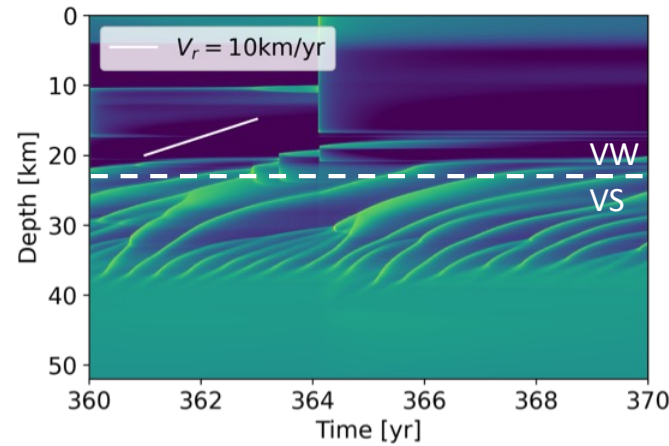
(a) Model A



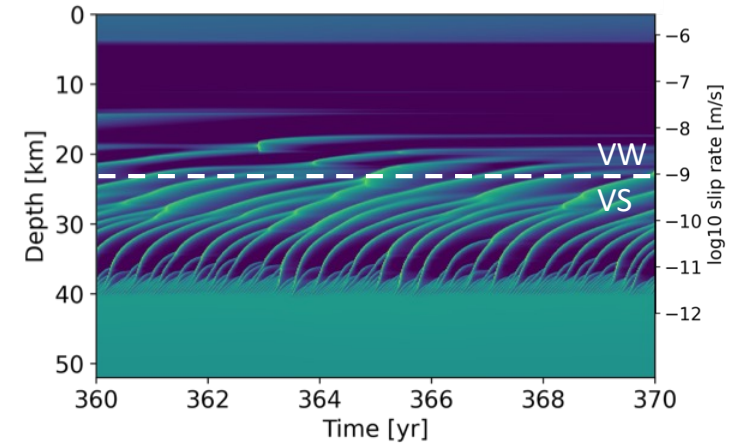
Reducing effective stress to ~1-10 MPa in slow slip region (by reducing shallow permeability) increases recurrence interval and decreases slip/event



(a) Model A

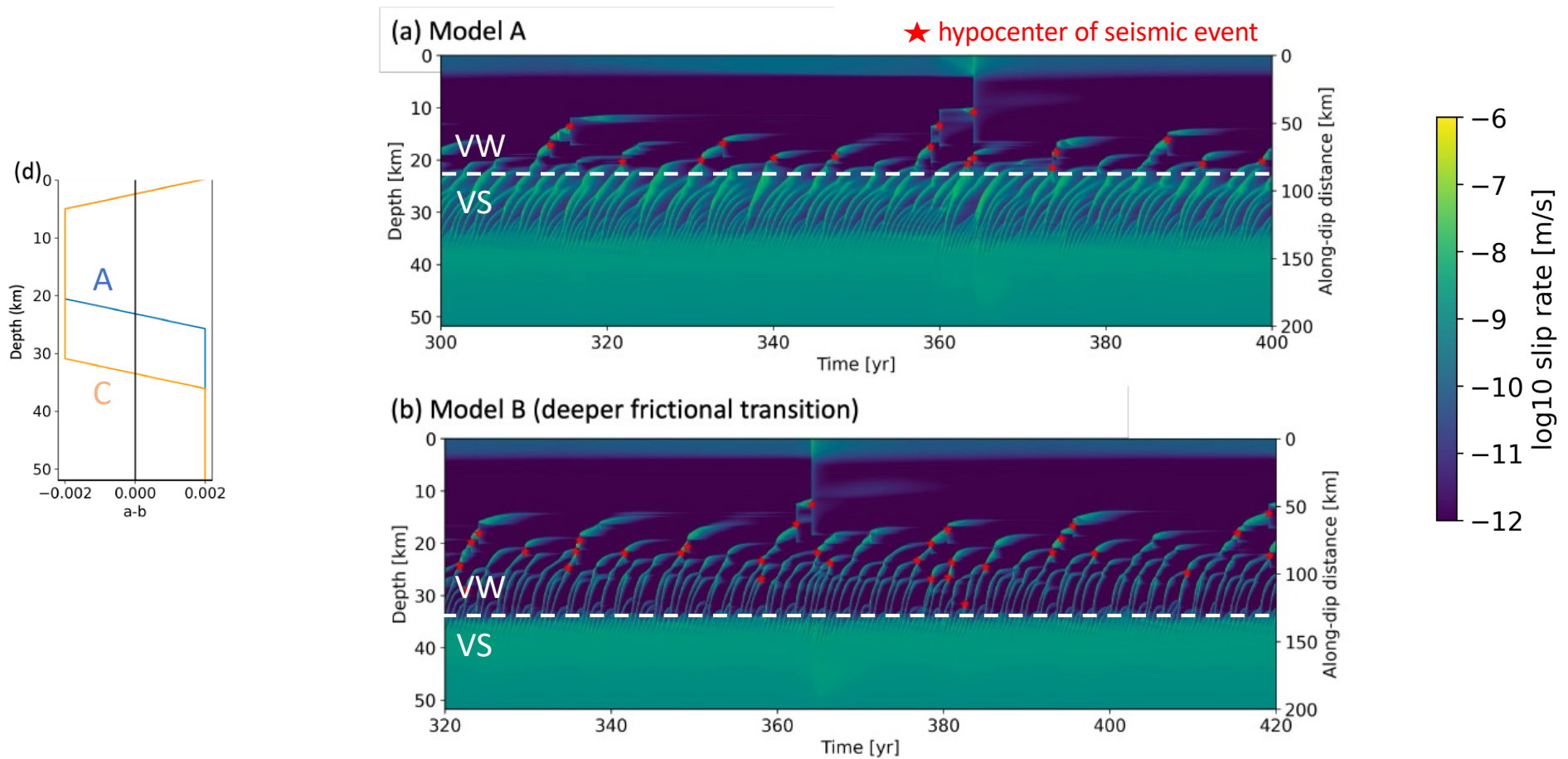


(c) Model C (low permeability)



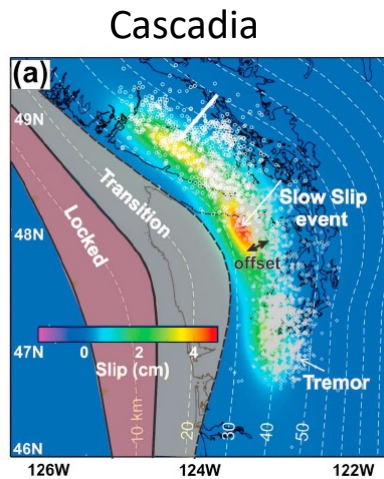
affects inter-SSE aseismic creep
and SSE propagation velocity,
recurrence interval, slip/event

Slow slip events “sharpen” if VW friction extends deeper

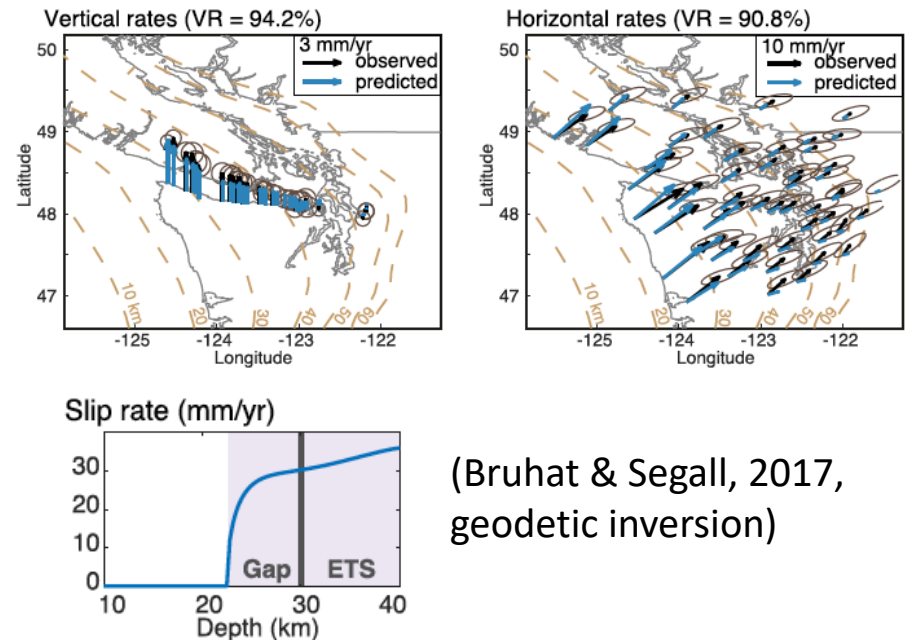


(Ozawa, Yang, & Dunham, 2024, in review)

Transition zone (or gap) between seismogenic zone and episodic tremor and slip (ETS) in some subduction zones



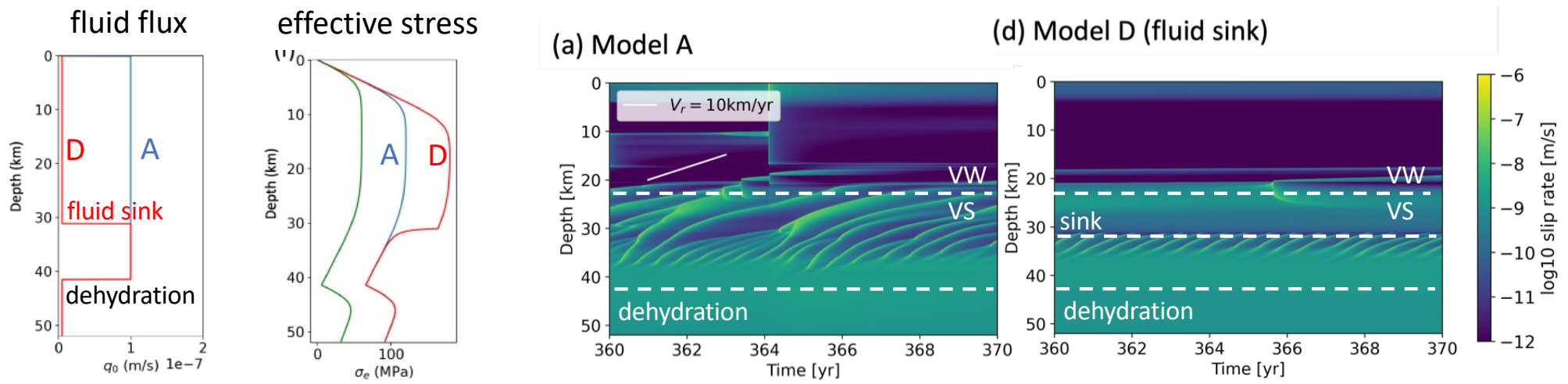
(Hyndman et al., 2015, after Kao et al., 2009; Wang et al., 2008)



(Bruhat & Segall, 2017, geodetic inversion)

Transition zone in Cascadia appears to be partially coupled (over past ~20 yr). Does it catch up in megathrust events? Or through some form of aseismic slip or distributed viscous shear? Major differences in hazard!

Fluid sink (near mantle wedge) corner creates transition zone between seismogenic zone and slow slip events



- sliding in transition zone is stabilized against fault valving by reduced flow rate
- slow slip events less complex, more periodic
- variable slip rate (coupling) in transition zone over cycle

Conclusions and future directions

Fault valving instability might explain slow slip

- Requires pressure-sensitive frictional slip, not distributed viscous flow
- Robust to frictional parameters (velocity-weakening or velocity-strengthening) and effective stress
- Recurrence interval (and slip/event) controlled by depth-dependent healing time
- Complexity (or periodicity) controlled by fault width (relative to length scale that depends on flow rate)

But:

- Propagation rates slower than observed (but increase as effective stress decreases)
- Propagation only in direction of fluid flow (for velocity-strengthening friction)
- 3D simulations (in progress!) required to understand along-strike propagation

Priorities for future work:

- *Fault-normal flow*, especially upward flow into overriding plate in subduction zones
- Accounting for *fault zone structure*, distinguishing between pressure & transport properties on slip surfaces (controlling fault strength) and in damage zone (controlling along-fault transport)
- Process-specific *porosity and permeability evolution models* (e.g., cracking & dilatancy in damage zone, vein formation & sealing by precipitation, silica transport)
- *Chemistry*: dehydration reactions; reaction kinetics for dissolution & precipitation of silica and similar minerals; lithology-dependent silica sources
- Coupling to *evolving temperature* from conduction and advective heat transport, as well as shear heating